

There has been much discussion in recent years about the effect of increasing global temperatures on marine fauna (see also the last issue of this magazine). However, it has often been overlooked that the increasing acidity of the oceans may have an even greater, and more insidious, effect on marine life than just a simple rise in temperature. Although acid rain, which contains nitric acid and sulphuric acid, falls into the oceans, it is the absorption of CO₂ from the atmosphere that is the most important contributor to oceanic acidity. This is because the shells of many marine creatures are made of calcium carbonate, which is a salt of carbonic acid, and which itself derives from CO₂. Any change, therefore, in the concentration of CO_a in sea water could have drastic consequences for marine invertebrates in particular and thereby marine life in general.

Acidification

Effect of absorbed CO, on marine fauna

Text by Michael Symes



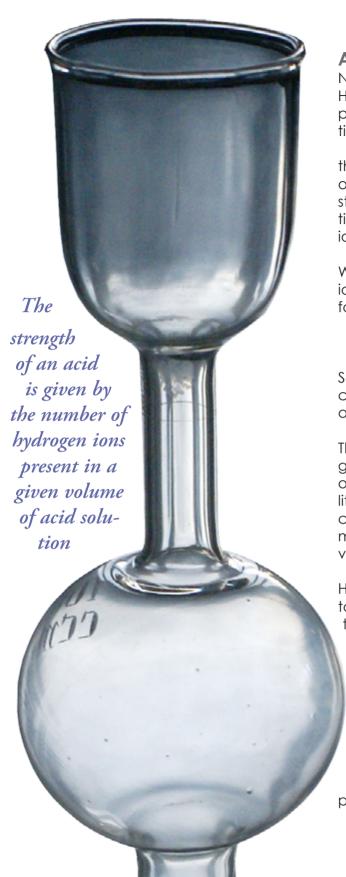
The absorbtion of CO₂ may have slowed global warming, but there has been a resulting change in seawater chemistry. To understand why the effect of increasing atmospheric CO₂, which leads to increased absorbtion of CO₂ in sea water, is important, it is necessary to understand some of the chemistry of carbonic acid and its salts. However, before we can discuss the effect of this acidity, we must have a useful description of what we mean by acidity and acid strength. The most common and useful unit to describe the magnitude of acidity is the pH value.

Measurement of acidity using the pH scale

When compounds, be they acids, alkalies or salts of these, dissolve in water they dissociate, wholly or partly, and in some cases extremely slightly, into their component ions.

For example, the gas hydrogen chloride, HCl, dissolves in water to give the ions H⁺ and Cl⁻. We call this solution *hydrochloric* acid.

Similarly, when the salt calcium sulphate, CaSO₄, dissolves in water the ions Ca⁺⁺ and SO₄⁻⁻ are formed.



Acid strength

Now, the strength of an acid such as HCl is given by number of hydrogen ions present in a given volume of acid solution.

If the acid is very concentrated and therefore very strong, the concentration of hydrogen ions will be very high. The strength of the acid is therefore proportional to the concentration of hydrogen ions.

Water partially dissociates into hydrogen ions and hydroxyl ions according to the following equilibrium:

$$H_2O <=> H^+ + OH^-$$

So water, like many other substances, can act both as an acid (having H⁺ ions) and as a base (having OH-ions).

The concentration of hydrogen ions, given in moles, is generally of the order of between 10⁻¹ and 10⁻¹⁴ moles per litre, with water having a hydrogen ion concentration of 10⁻⁷ moles per litre. This means that water dissociates to only a very small degree.

However, these are cumbersome values to use, so the hydrogen ion concentrations are converted to the much more convenient pH values.

This is done by taking the negative of the log¹⁰ of the hydrogen ion concentration. To give an example—If the hydrogen ion concentration is 3x10⁻⁶ moles per litre then the pH value is given by:

$$pH = -\log_{10}(3 \times 10^{-6})$$

 $= -(\log_{10} 3 - 6)$

= -(0.48 - 6)

= 5.52.

It will be seen, then, that a decrease of one unit in the potential of the hydrogen ion concentration, say from 10-4 to 10⁻³, means a decrease of one pH unit.

To understand why the effect of increasing atmospheric CO which leads to increased absorbtion of CO, in sea water, is important, it is necessary to understand some of the chemistry of carbonic acid and its salts.

Accordingly, a unit decrease in pH value means a ten times increase in the hydrogen ion concentration i.e. a ten times increase in acidic strength.

The molality of hydrogen ions in pure water is actually 1.004x10⁻⁷ mol kg⁻¹. The pH value of pure water is therefore -0.0017 + 7 i.e. 6.998 or to 2 decimal places 7.00.

As water is a neutral substance from the acid/base point of view, a pH value of 7 is considered as being midway between acid and alkali.

An aqueous solution of HCl of unit molarity, i.e. 1 mole per kg, with virtually complete dissociation, will have a pH of 0.09.

Some typical pH values

Battery acid: ca 1 Gastric juices: 1 – 3 Lemon juice: ca 2.2

Household vinegar: 2.5 to 3

Tomato: 4.0

Cheddar cheese: ca 5.90 Camembert cheese: ca 7.44 Concentrated solution of baking

soda: 8.2

Household cleaning fluids: 10 to 11



As stated above, the acidity of seawater mainly arises from the amount of CO₂ in it and the acids it forms. Carbon dioxide is highly soluble in water in which it forms an equilibrium with carbonic acid,

About 0.5% of the dissolved CO₂ remains as H₂CO₂ with the rest immediately dissociating in the two steps given by the following equilibria:

Firstly,

 $H_2CO_3 \iff H^+ + HCO_3^-$, forming the bicarbonate ion.

Then,

HCO3 <=> H+ + CO3 , forming the carbonate ion.

In absolute terms there are, on average, about 45ml total CO, per litre of sea water, but because of these two equilibrium reactions, nearly all of this is found in the carbonate and bicarbonate ions. Only about 0.23ml CO, per litre occurs as dissolved gas in the water i.e. 0.51%.

Hydrogen ions are liberated in these equilibrium reactions, which means that the pH of seawater is largely determined by the concentra-

tion of bicarbonate and carbonate ions. The pH is usually 8 ± 0.5 . However, when CO₂ is added to seawater, more H_oCO_o is formed and the number of

Calcium carbonate is an important constituent of coral reefs and mollusc shells

hydrogen ions increases. The pH thereby decreases, with the seawater thus becoming more acidic—though it might be more descriptive to say less alkaline.

The carbonate ion can interact with cations such as Ca⁺⁺ and Mg⁺⁺ in the seawater causing them to precipitate as insoluble carbonates. Calcium carbonate is the main constituent of chalk, limestone and marble. It has a solubility of only about 20mg per litre of cold water, and because it is so insoluble, it is an important constituent of coral reefs and mollusc shells.

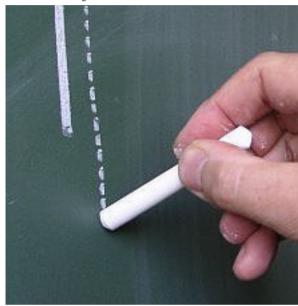
Note: It is usually claimed that the pH scale runs from 0 to 14. However, there is no fundamental reason why pH values cannot exist outside this range. For example, a 2 mole per kg solution of HCl in water has a pH value of -0.31. Also there is no reason why a very strong basic solution should not have a pH value exceeding 14. For most purposes, however, it is sufficient to confine our attention to the range 0 to 14.



EQUIPMENT BOOKS SCIENCE & ECOLOGY

However, and this is important for understanding the effect of increasing CO. in sea water, this calcium carbonate, although it is insoluble in water, dissolves in acidic solutions.

For example, as many of you will have seen, the strong acid hydrochloric acid, HCI, will dissolve chalk. This gives a solution of calcium chloride and carbonic acid, which then dissociates to produce water and the emission of a froth of bubbles of CO₂.



Acids will dissolve chalk. Chalk is made by minute calcite plates shed from micro-organisms called coccolithophores

This is shown in the following reaction. Note that it is not an equilibrium; it only goes one way as the CO₂ is continually removed as a gas.

$$CaCO_3 + 2H^+ + 2CI^- \rightarrow Ca^{++} + 2CI^- + CO^2 + H_2O$$



When acids dissolve, chalk CO₂ is released

JONATHAN ZANDER Calcium carbonate is an important constituent of coral reefs and mollusc shells

With carbonic acid, however, which is a weak acid, a different reaction occurs. When there is excess CO₂ dissolved in the water, the following equilibrium occurs in which the bicarbonate ion is formed:

Equilibrium A

$$CO_{2} + H_{2}O + CaCO_{3}$$
 <=> $Ca^{++} + 2HCO_{3}^{-}$

The right hand side of the equilibrium is thus a solution of calcium bicarbonate that, unlike calcium carbonate, is quite soluble, having a solubility of 166 g per litre. However, it exists only in solution. This solution contains calcium ions, carbon dioxide, bicarbonate ions and carbonate ions.

As this is an equilibrium, the more CO₂ pushed into the left hand side of the equilibrium, the more it will force the equilibrium to the right i.e. the more the insoluble calcium carbonate will be changed into the soluble bicarbonate.

Incidentally, the bicarbonate ion is useful in the kitchen were it is to be found in NaHCO₃, sodium bicarbonate, better known as baking soda. When it combines with moisture and an acidic ingredient such as milk (pH about 6.4), bubbles of CO_o are formed that expand under the oven temperature, causing the cake, or what ever is being made, to rise.

Reaction B is thus important for understanding certain geological processes. However, it is reaction A that is the most important for understanding the effect of CO₂ on life on Earth. It is fundamental, in fact, for understanding the effect of the increase of CO₂ on the chemistry of seawater and on the life in it.

Unlike sodium bicarbonate, which exists

as a dry solid, calcium bicarbonate does

not. If an attempt is made to prepare

lowing reaction.

Reaction B

calcium bicarbonate by evaporating

its solution to dryness, solid cal-

cium carbonate will be obtained

and CO₂ as shown in the fol-

aqueous Ca(HCO_3)₂ \rightarrow

This reaction is important, for example, in the formation of speleothems such as stalactites and stalagmites.

Water picks up CO₂ as it passes

through the atmosphere and through the topsoil. As it flows past limestone, chalk and other minerals containing calcium carbonate, it dissolves some of this carbonate, forming bicarbonate

by the process given in equilibrium A

above. When this solution reaches the roof of a subterranean cave, the water

will evaporate, and insoluble calcium

carbonate will be formed as shown by

tially where the evaporation is greatest

i.e. at the lower ends of protuberances,

reaction B. As this is formed preferen-

increasingly longer formations will be

instead, together with water

gaseous CO₂ +

gaseous H₂O +

solid CaCO

Equilibrium A is fundamental for understanding the effect of the increase of CO2 on the chemistry of seawater and on the life in it

formed.



Acidity of the oceans The acidity of the oceans can arise from two main sources.

- · Acid rain, which is formed when rain water absorbes CO_a, together with nitric acid and sulphuric acid, on its fall through the atmosphere.
- Direct absorption of CO₂ in the surface water layer.

Although acid rain is obviously important for terrestrial life, it seems to play a minor role in making sea water more acidic on a global scale. Its impact, though, is somewhat areater in shallower waters such as the coastal regions. It is the direct absorbtion of CO_a by the surface of the sea water itself and the corresponding impact on its chemistry that affects marine creatures. We therefore concentrate here on the acidity arising from absorption of CO₂.

Since about 1800 i.e. from the start of industrialisation in Europe, the amount of CO, in the atmosphere has increased by about 36%, from 280 parts per million (ppm) to 380 ppm i.e. 0.038%. It is claimed that nearly 50% of the CO₂ released due to human activities over the last 200 years has been absorbed by the oceans.

Currently, the oceans absorb about 22 tons of CO₂ per day. However, the oceans tend to mix rather slowly and, because CO₂ mainly enters the water by gas exchange at the surface, most CO₂ is found in the top layers of

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the oceans or in the shallow seas. This means that the acidity is not completely uniform throughout the oceans. Deep cold waters, which only circulate to the surface very slowly, are far from being saturated. In 2005, the warm waters around the widely separated islands of the Canaries, Hawaii and Bermuda, had a pH of ca 8.09.

It is important to remember, especially in these times of global warming, that warm water dissolves less CO_a than cold water. This means that, according to equilibrium A, this will favour the production of carbonate over bicarbonate.

As stated, the pH of seawater varies both in time and place, and due to the equilibrium A given above, even small local changes in pH towards the acidic can have great effects on marine life. For example, as the deep cur-

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rents circulate around the globe, they eventually arrive in the North Pacific where they rise to the surface before plunging down deep again to continue on their deep journey. As it upwells, the cold water absorbs more CO₂ from the atmosphere, which means that equilibrium is pushed to the right giving an increased production of bicarbonate ions at the expense of carbonate ions.

Effect on marine fauna

Most of shell-forming molluscs which utilise CaCO, to form their shells belong to two classes: the Gastropoda i.e. the univalves, snails, and the Bivalvia i.e. the clams, oysters, scallops, etc. Other classes of molluscs that form calcium carbonate shells are the Scaphopoda (the tusk shells), the Polyplacophora (the chitons) and the Monoplacophors (for example the deep sea Neopilina). Some echinoderms such as starfish, sea urchins and sand dollars have hard exoskeletons formed of calcium carbonate, as do some annelid worms; and so, of course, do the reef structures of corals.

All shells of the mollusca are composed of basically the same material. This is calcium carbon-

ate, in the form of calcite or aragonite, crystallised out in an organic matrix and laid down in layers by the shell alands of the mantle tissue. Aragonite and calcite are two polymorphs of CaCO, in which the crystal structures differ. Aragonite forms naturally in almost

all mollusc shells as well as in the endoskeletons of both warm- and cold-water corals. In some mollusca, the entire shell is aragonite, while in others, it is a mixture of both aragonite and calcite. These shell-forming animals produce their calcium carbonate shells or skeletons from the CO₂-- ions in the sea water.

The shells of crustacea, such as lobsters, crabs and shrimps are different in that they are built up from only 30-50% calcium carbonate together with 30-40% protein and 20-30% chitin. Chitin is a polysaccaride polymer also found in the exoskeletons of terrestrial insects.

Now, we have seen that the insolubility of calcium carbonate in seawater depends

As the concentration of CO, in the water increases, the carbonate ion concentration decreases. There will therefore be less material for animals to build their shells

on its acidity i.e. its pH value; and as the concentration of CO_a in the water increases, the carbonate ion concentration decreases. There will therefore be less material for animals to build their shells. This will inhibit calcification i.e. the process by which these animals rebuild

their shells, and the skeletal arowth rates of calcareous plankton, for example, will be reduced. Not only that, the increasing acidity can actually dissolve away the shells of oysters, clams, krill and pteropods, so that in areas of high CO_a concentration these animal will die. Creatures with shells living at higher, colder, latitudes, and also in near-shore waters, will therefore probably be in the most trouble. For example, shell-building planktonic organisms such as the coccolithophorids and foraminifera are an important basis of the food chain in the waters of Antartica which.



Iincreasing acidity can actually dissolve away the shells of oysters, clams, krill and pteropods





being cold, can absorb greater amounts of CO₂. The increasing CO_o content of sea water will also affect the thickness and strength of lobster and other crustacea shells thus affecting their ability to survive in their normal habitats.

Although protected by their spines, echinoderms are preved upon by, among others, the triton shell, the trigger fish, crabs and shrimps, and by other carnivorous echinoderms. Echinoderms also serve as hosts to many symbiotic shrimps, crabs, worms, snails and even some fish. Any change in the ability of these creatures to form their protective armour, and

to survive in their present environment, could therefore be a disaster for both host and prey.

Krill and pteropods are major food source for iuvenile salmon, herring, pollock,

cod mackerel and many other fish. Disturbing the lower end of a food chain can have dramatic effects at the higher end of the chain. As these organisms provide food and habitat for other species fewer of these animals will then have a feedback effect on the populations of their predators and their prey. Entire eco-systems can therefore change or even completely disappear for ever.

Although unaffected by decalcification, squid are also sensitive to higher acidities, which affect their blood circulation and respiration.

The future

haps, but we must assume

that humans are sensi-

ble enough, and clever

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The situation is very complex. The increasing amount of CO₂ in the atmosphere will lead to an increased uptake of CO_a in the oceans. However, the increasing level of CO₂ is leading to an increase in global temperature

thereby causing It is somewhat naive per- ocean water temperatures to rise. This increase in water temperature will then decrease the solubility of CO_a. So, will we arrive back at the status quo or will there be a permanent change in the

CO. content of the oceans? At this time, it is not known how this interaction will be resolved.

Some things do already appear to be clear, though. The surface waters will become more and more saturated with CO₂ so that in the short term, the oceans will probably become a much less efficient sink for CO_a. This will obviously have an effect on the Earth's climate.

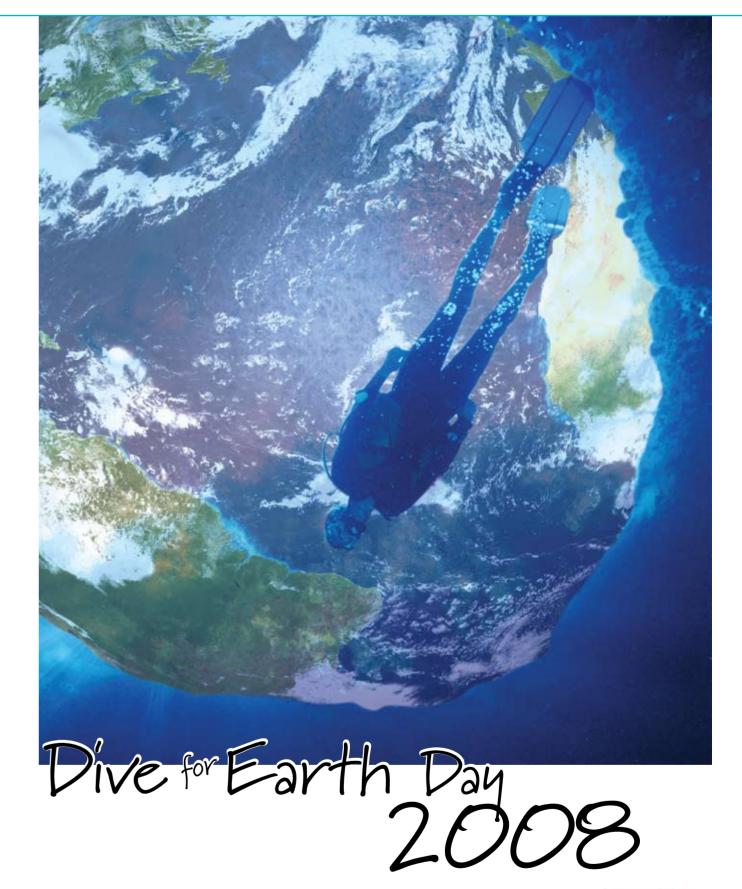
100 years, it is estimated that the pH of the oceans could drop by half a unit to about 7.7 i.e. a 5 times increase in acidity, with the amount of carbonate available to marine organisms dropping by 60%. And some scientists believe that in less than 50 years, the oceans could become too acidic for corals to survive at all.

SOURCE: COURS DE SCIENCES NATUREUES 1937

say, less than

Although the future may seem somewhat black for the marine fauna, the increase in CO₂ content of the oceans may actually be beneficial to a number of marine flora, for example the sea grasses, although this is not cer-

In the long run, it is estimated that over a period of several thousand years about 90% of the anthropogenic CO₂ emission will end up in the oceans, and that the pH of the oceans could be lowered permanently. It is somewhat naive perhaps, but we must assume that humans are sensible enough, and clever enough, to reduce CO₂ emissions to a level below that which they are today. What the effect on marine life will be of all the already accumulated CO₂ in the oceans is, however, unclear. One thing seems to be certain though: it probably won't be a benign effect. ■



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